

# **Solar variability and climate**

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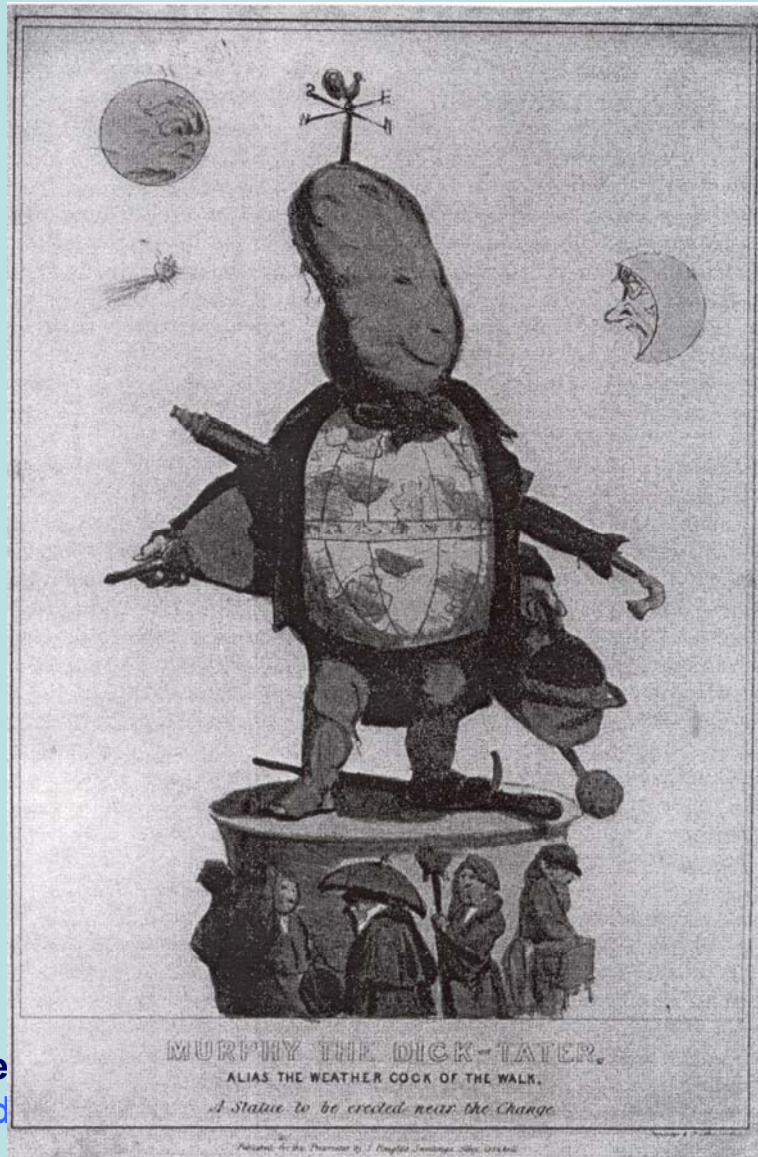


# Outline

- Why has this subject historically been viewed with scepticism?
- What are the observed impacts of solar variability on climate?
- How much does solar output vary – now and in the past?
- Can we explain the observed climate impacts?

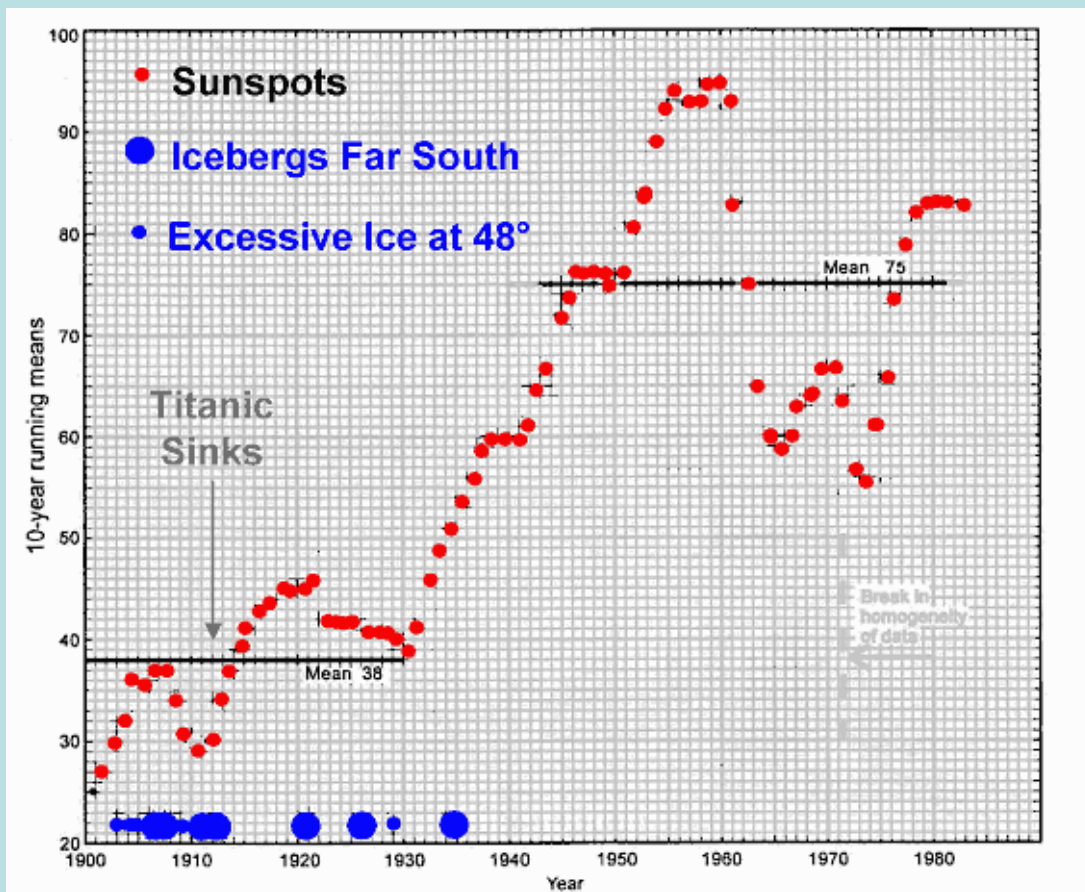
# Astro-meteorologists

Patrick Murphy  
(satirical engraving 1838)



Richard Morrison, a.k.a. Zadkiel  
(*Punch*, 1863)





Lawrence, *Weather*, 2000

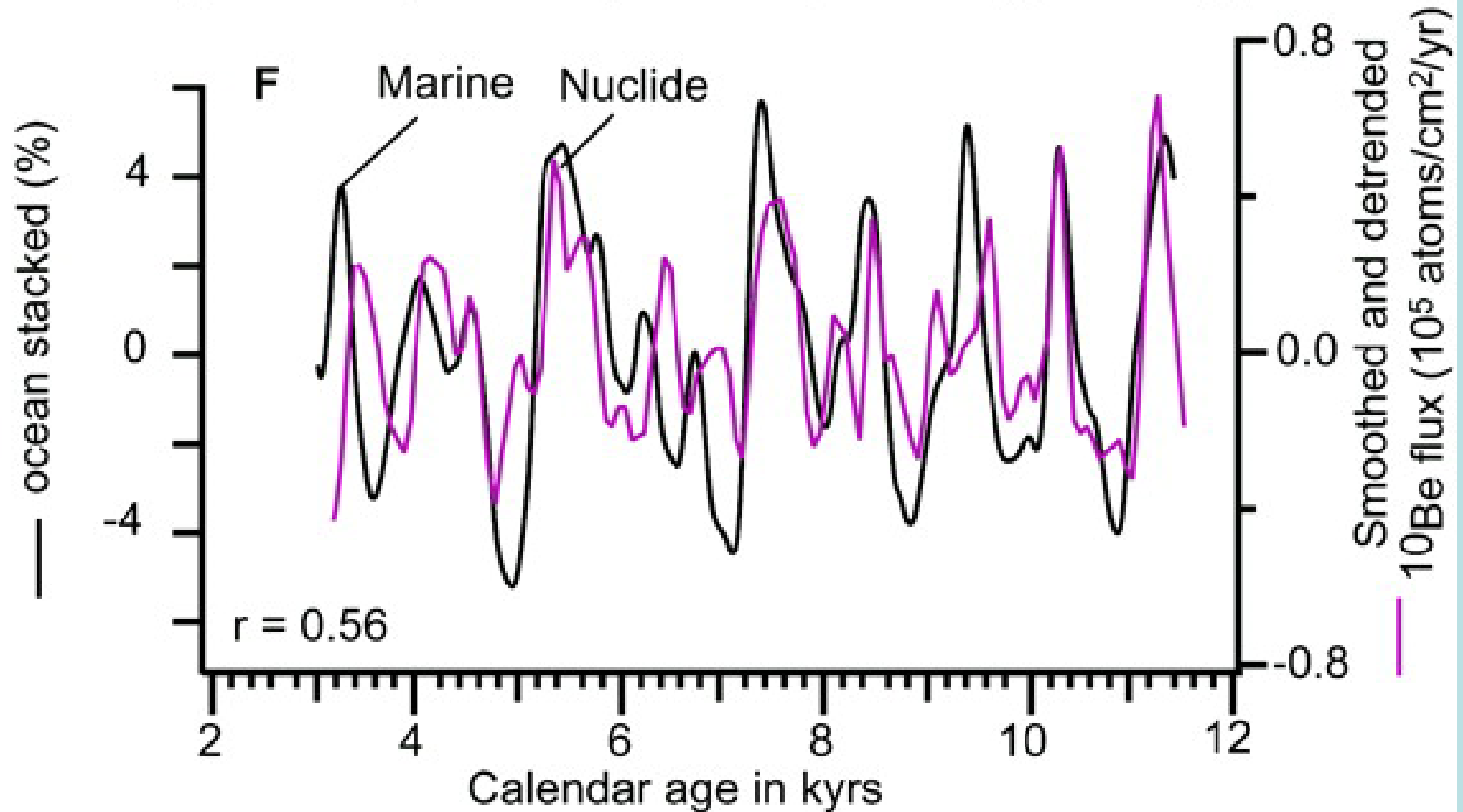
## The Sun-Climate Connection (Did Sunspots Sink the Titanic?)

*NOAA website, 2001*

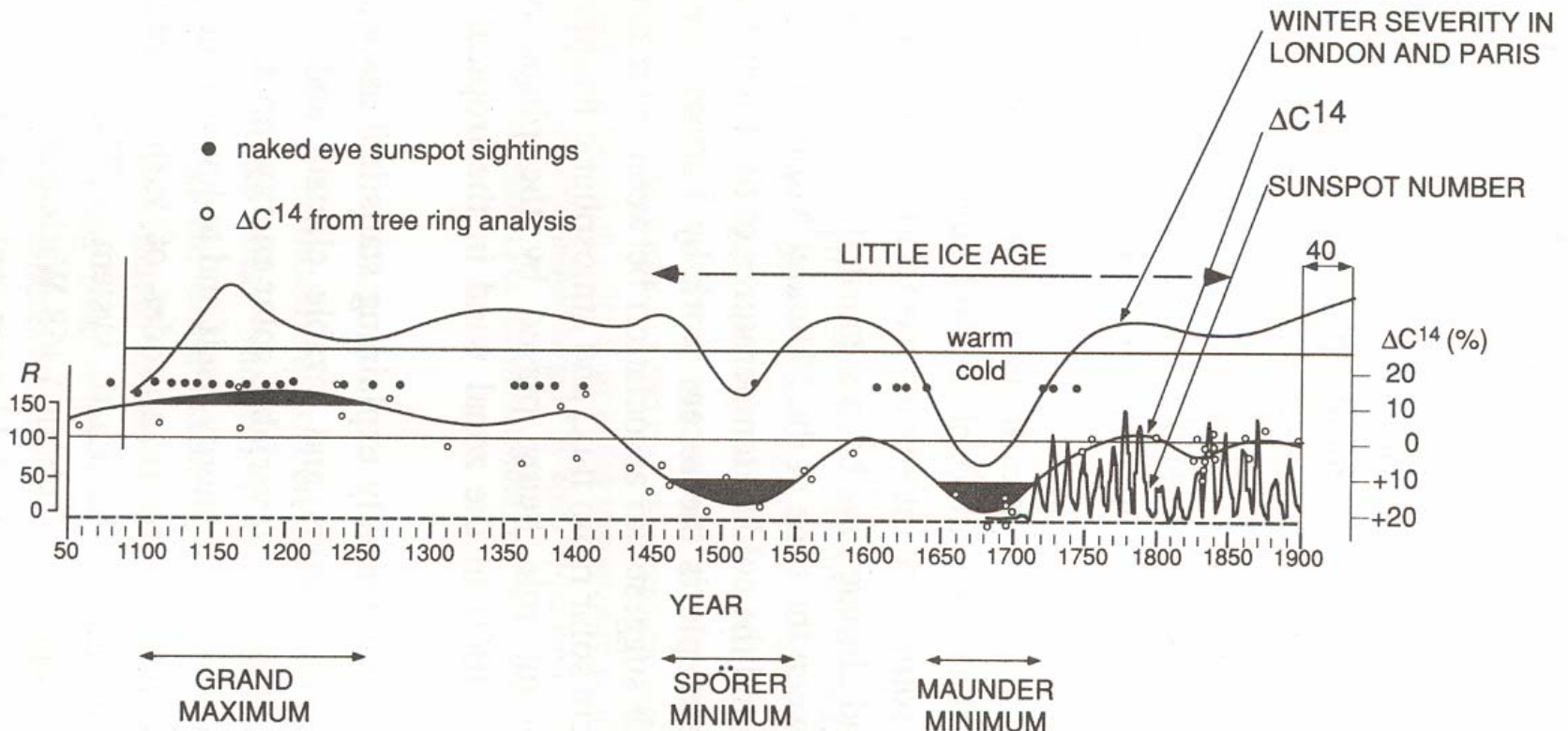


# N.Atlantic ice flow-deposited sediment

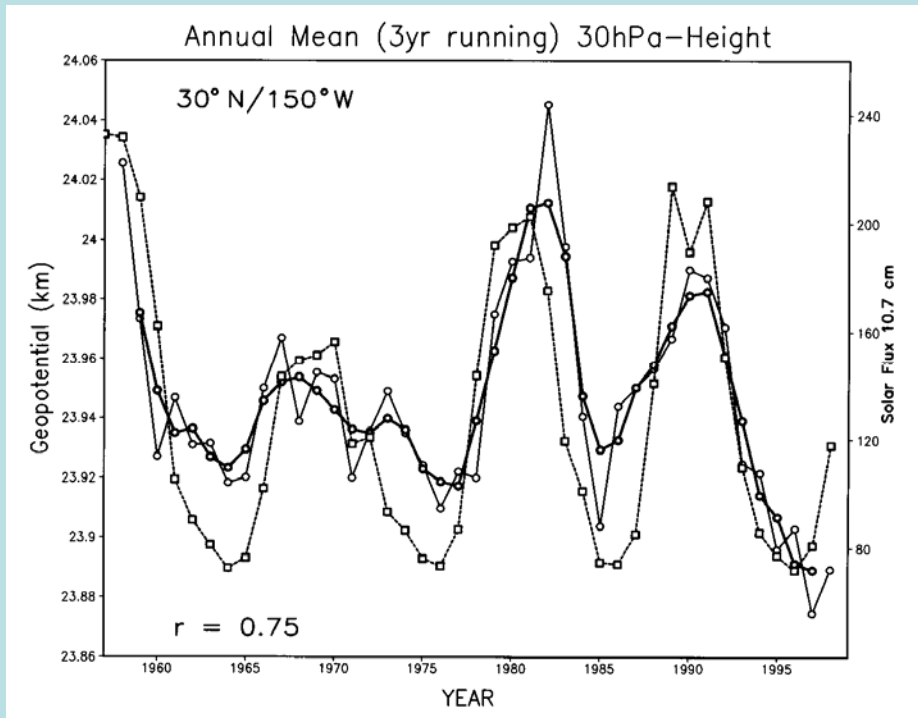
Bond et al, *Science*, 2002



# Temperature in NW Europe



## 30 hPa geopotential height (annual mean, Hawaii)

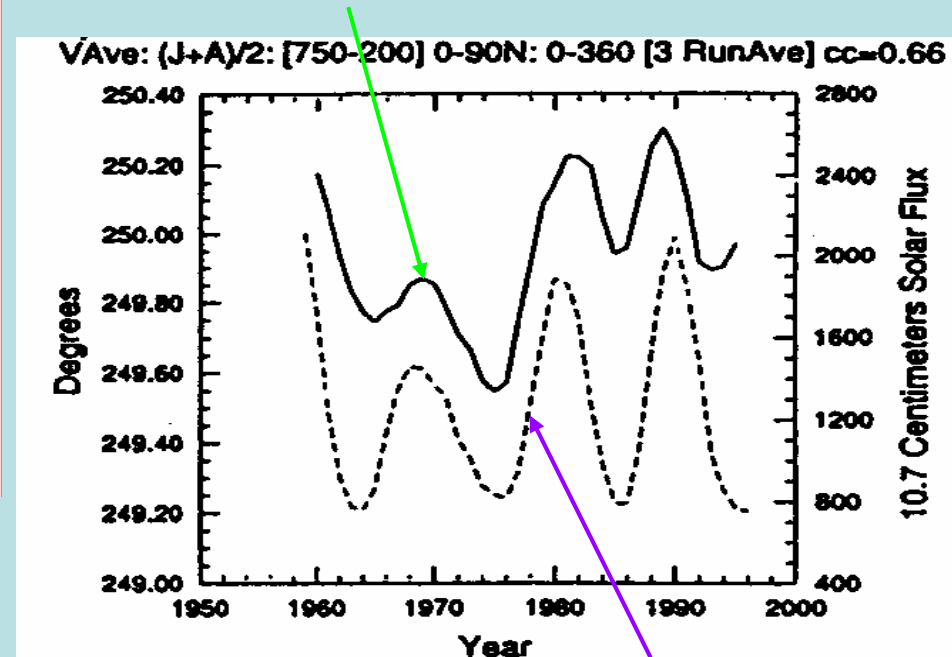


Labitzke and van Loon (1995)

Imperial College  
London

## Upper troposphere temperatures (NH summer)

750-200hPa thickness  
July & August  
3-year running average

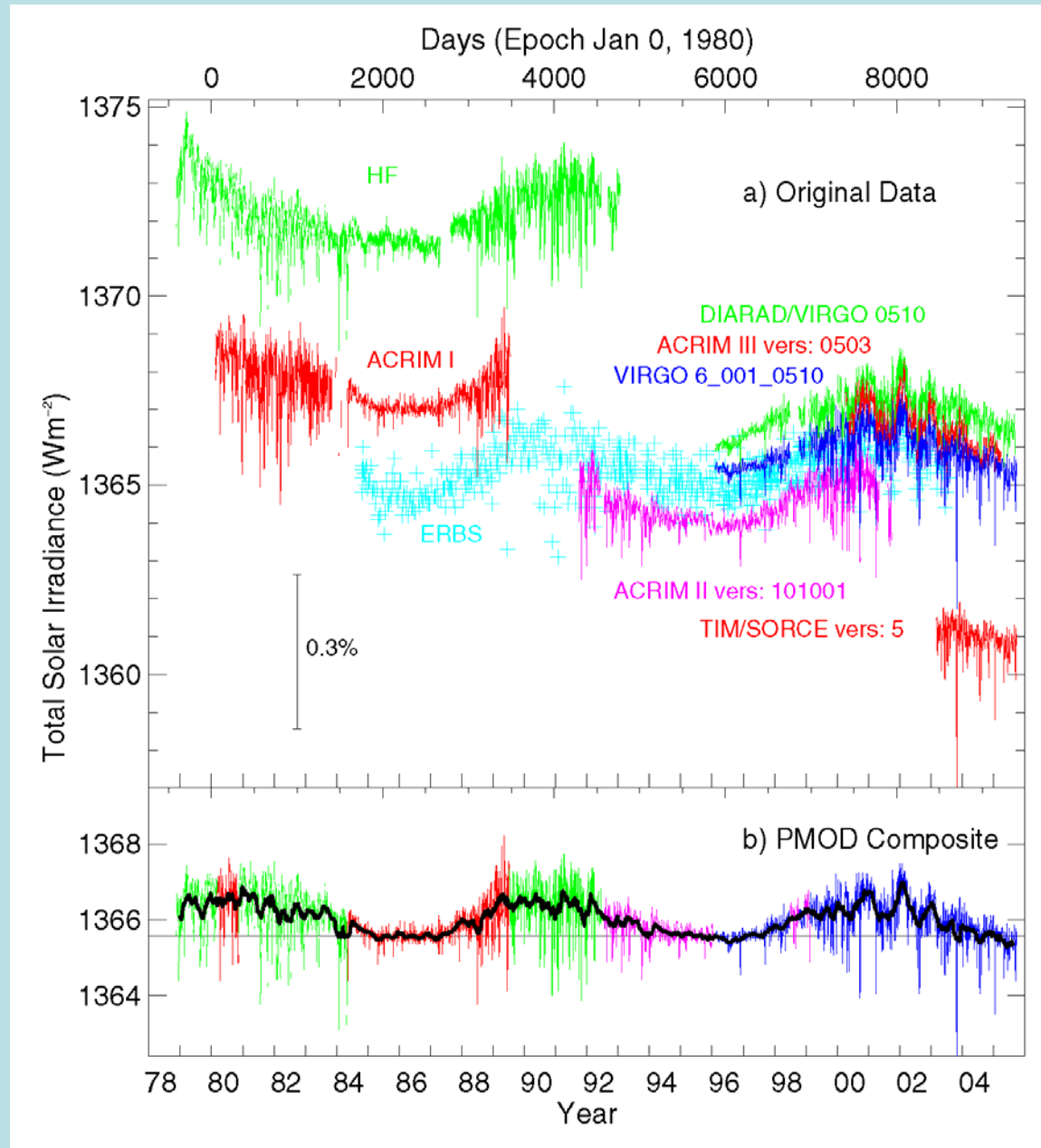


solar 10.7cm flux

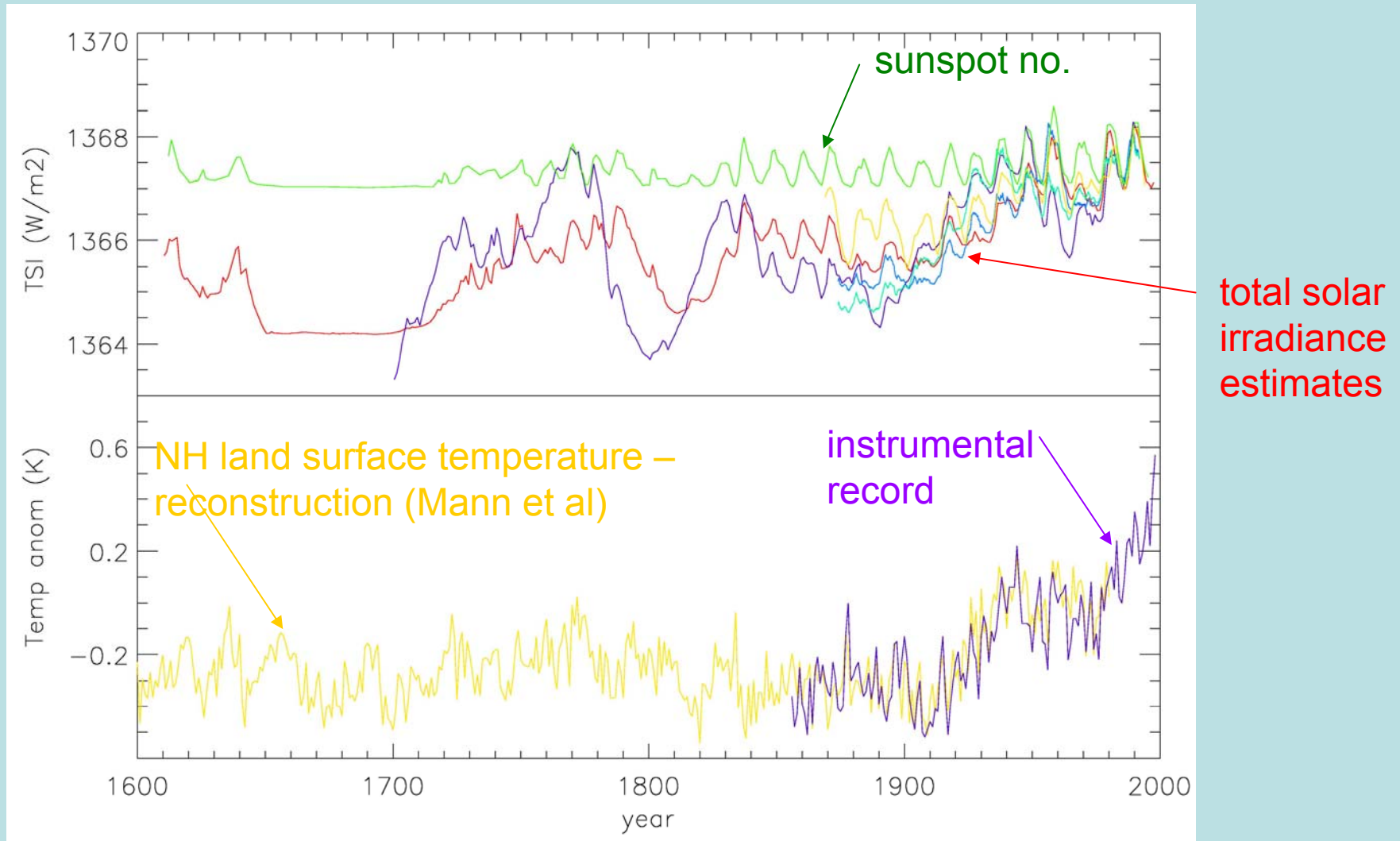
van Loon and Shea (2000)

# Total solar irradiance (TSI) measurements from satellites

courtesy Claus Fröhlich (PMOD)



# “Reconstructed” solar irradiance & NH land temp.



# Radiative forcing

Useful because GCMs, & limited observational studies, suggest that the perturbation to global average, equilibrium surface temperature,  $T_g$ , is related to radiative forcing,  $RF$ , by:

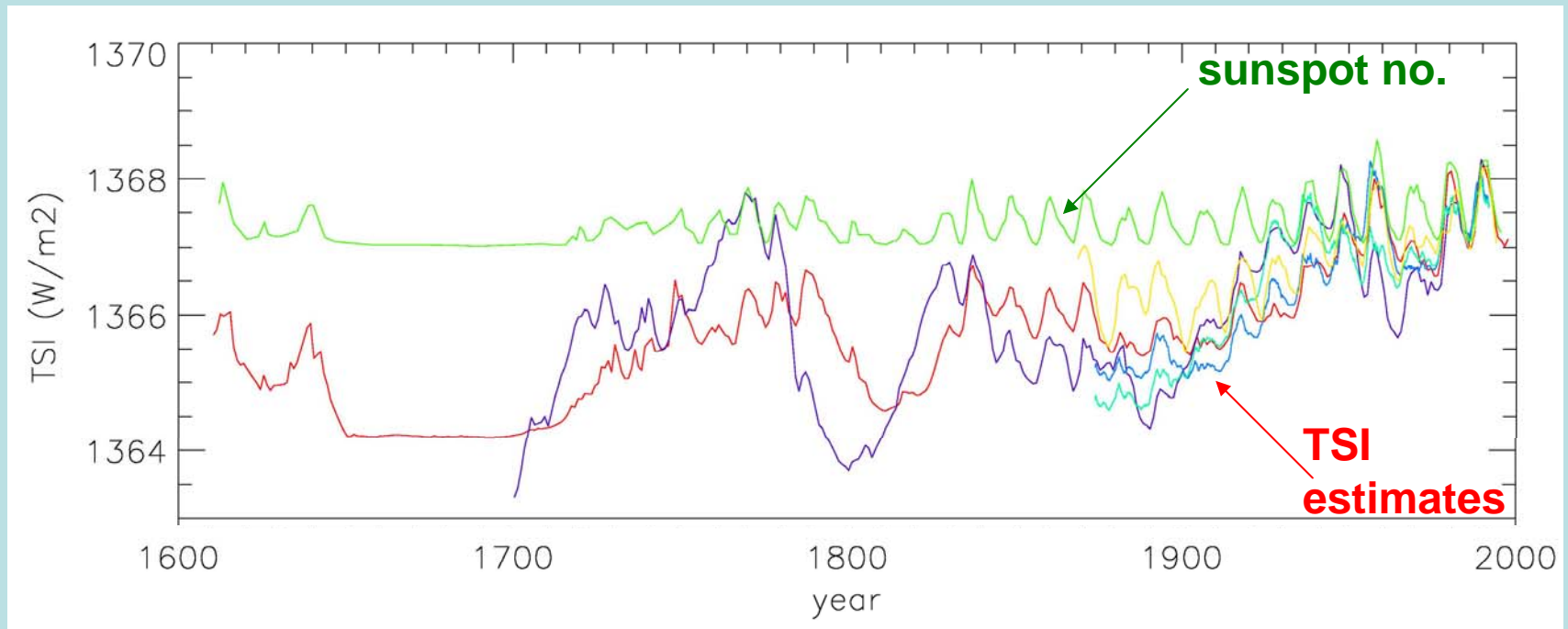
$$\Delta T_g = \lambda RF$$

where  $\lambda$ , the climate sensitivity parameter, is independent of the nature of the forcing.

$$\lambda \sim 0.6 \text{ K (W m}^{-2}\text{)}^{-1}$$

$$[ 0.3 < \lambda < 1.0 ]$$

# TSI and radiative forcing

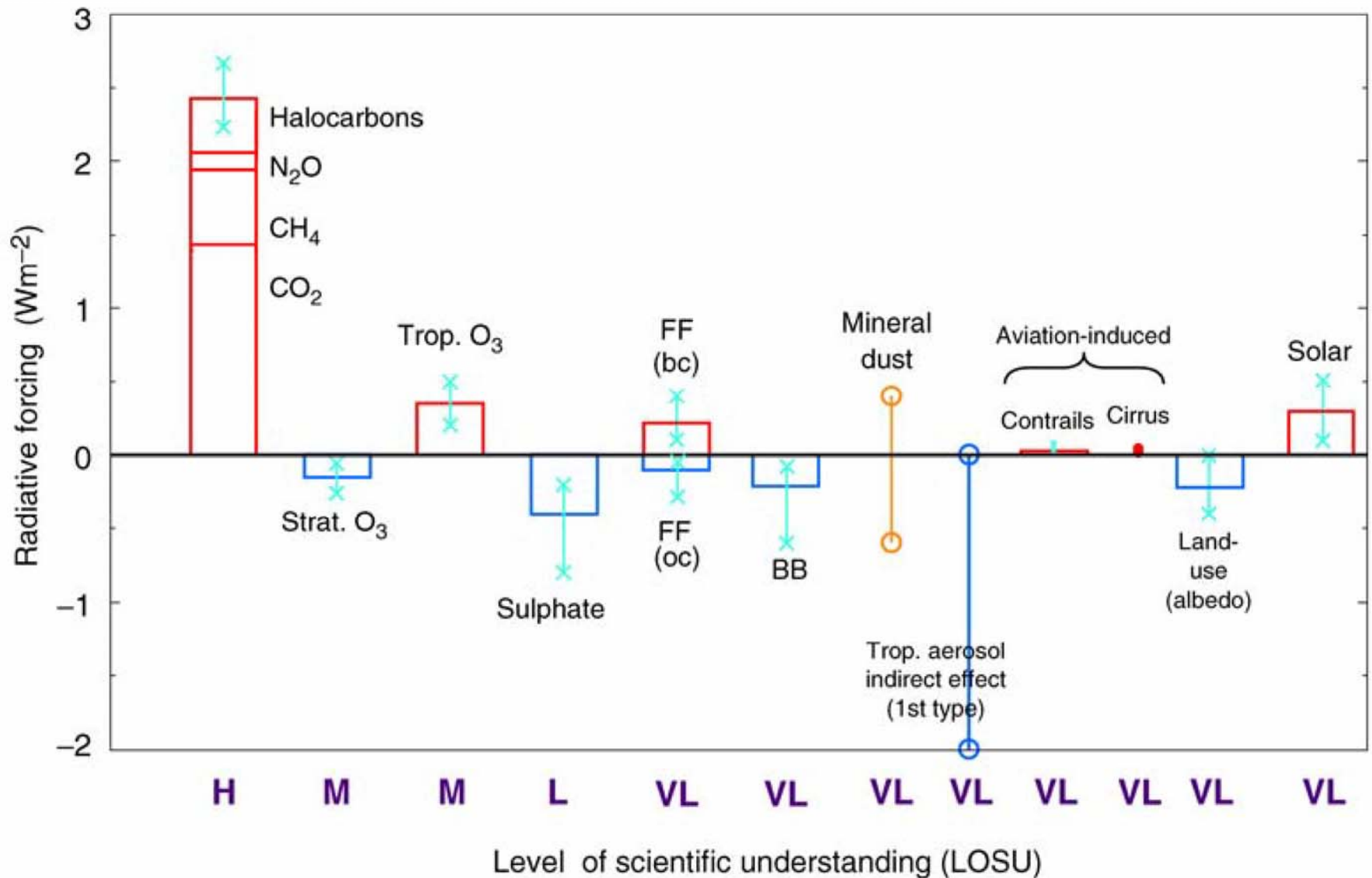


$$RF = \frac{1}{4}(1 - \alpha)\Delta TSI \approx \frac{\Delta TSI}{6}$$

$\Delta TSI$  of  $\sim 2.5 \text{ W m}^{-2}$  since 1900 implies  $\Delta T_g$  of  
 $\frac{2.5}{6} \times 0.6 = 0.25 \text{ K}$  cf  $\sim 0.7 \text{ K}$  in temperature record

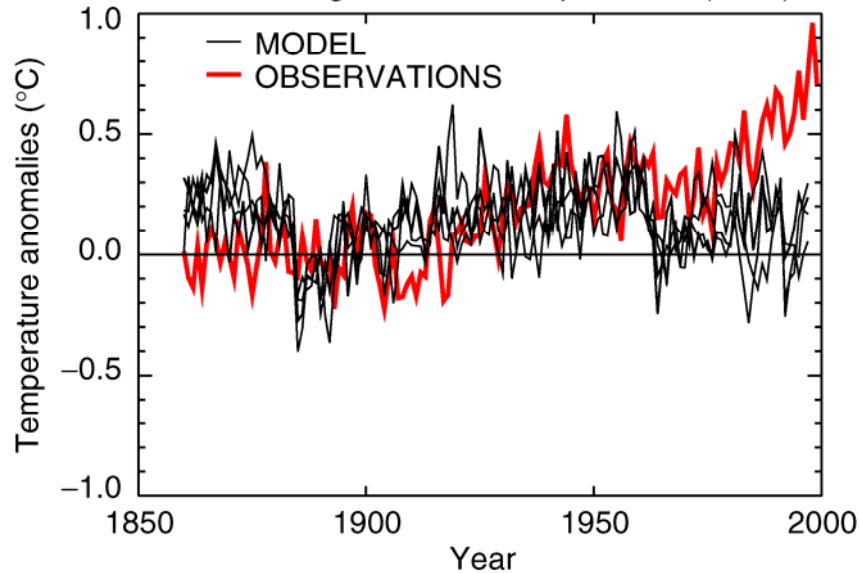
# Intergovernmental Panel on Climate Change (2001)

Global and annual mean radiative forcing (1750 to present)



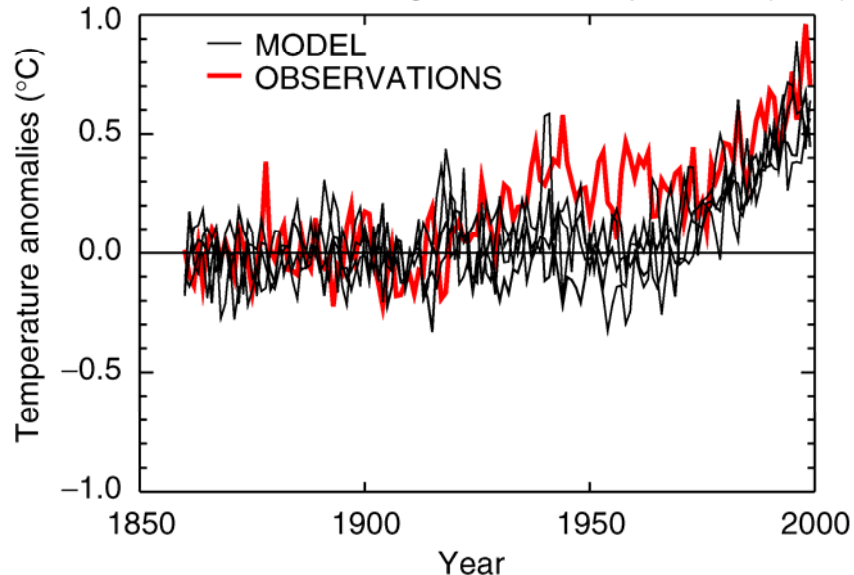
(a)

NATURAL : Annual global mean temperatures (1.5m)



(b)

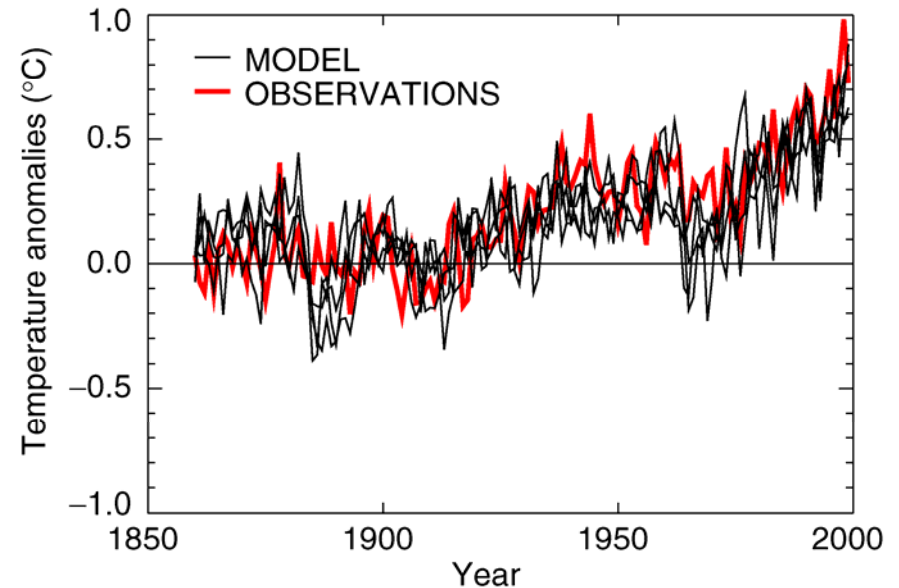
ANTHROPOGENIC : Annual global mean temperatures (1.5m)



# GCM global mean surface temperature simulations IPCC 2001

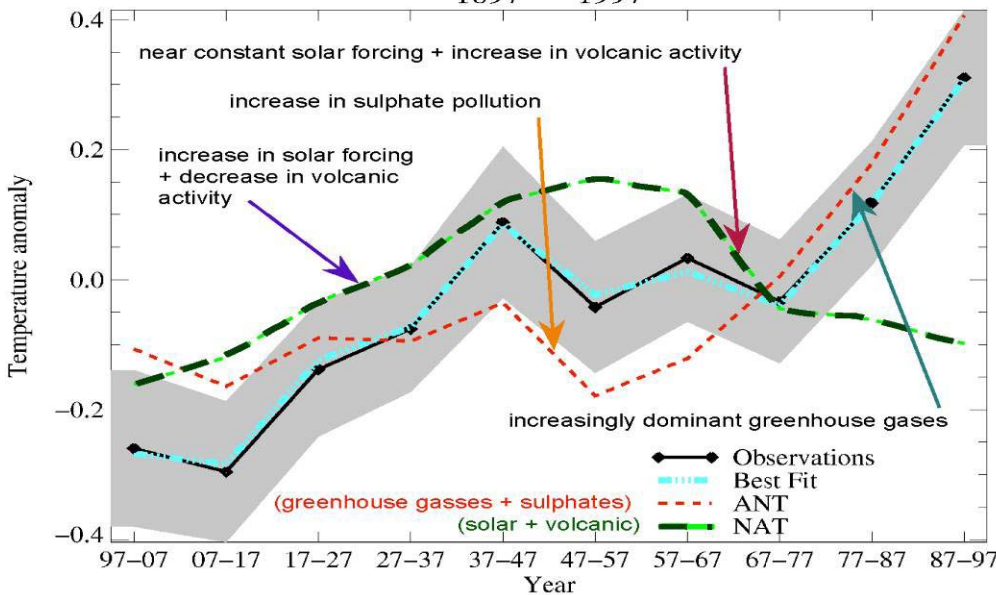
(c)

ALL FORCINGS : Annual global mean temperatures (1.5m)

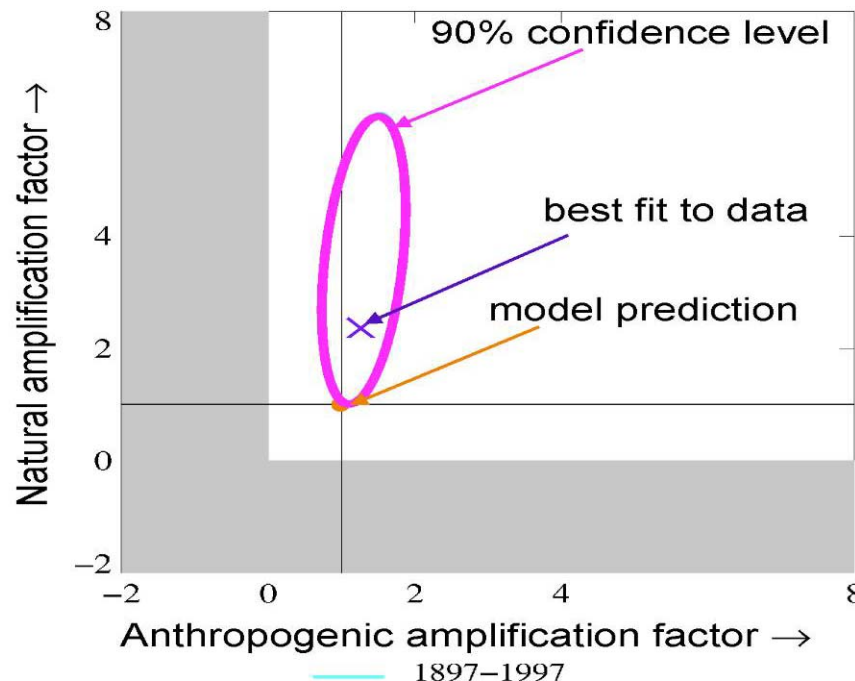


***So all is explained?***

1897 — 1997



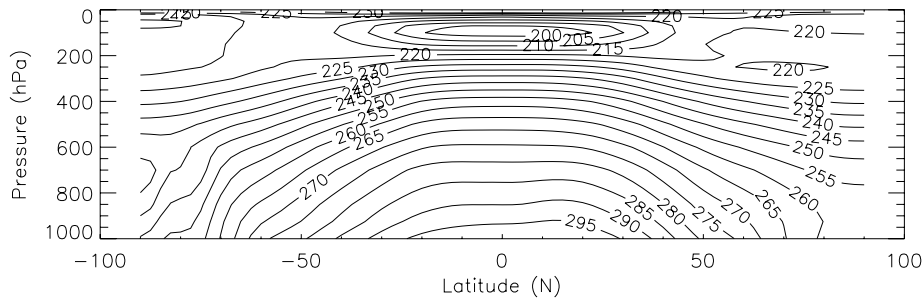
# Detection/attribution of surface temperature trends



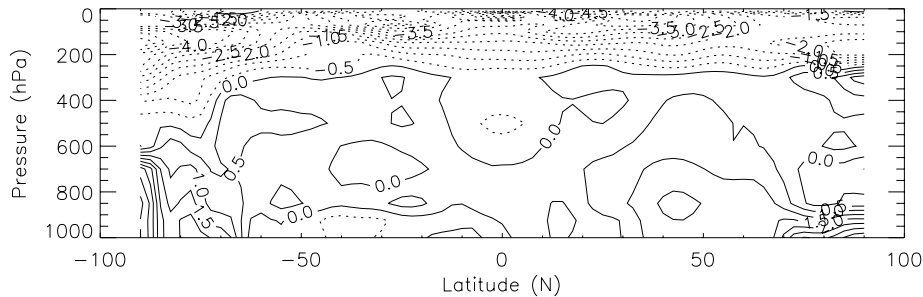
**No!**  
**Model underestimates response to natural forcing by a factor of 2 or 3**  
 Stott et al 2003

# Influences on temperature

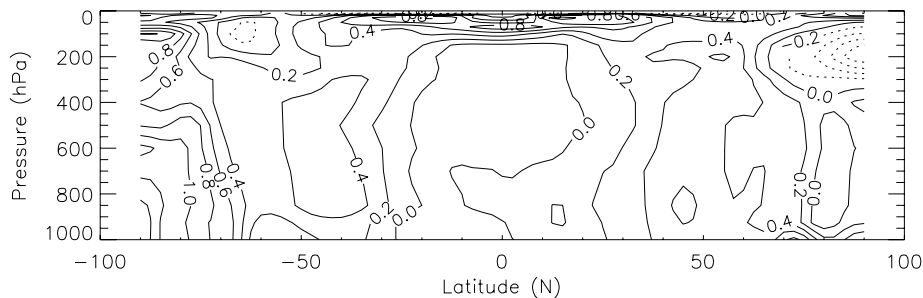
mean



trend

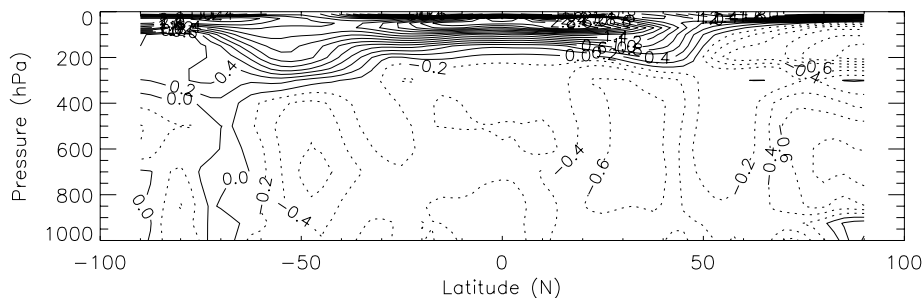


solar

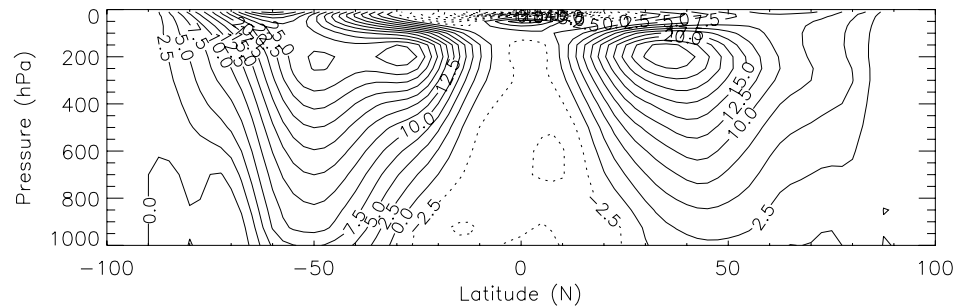


***Not a uniform warming!***

volcanic

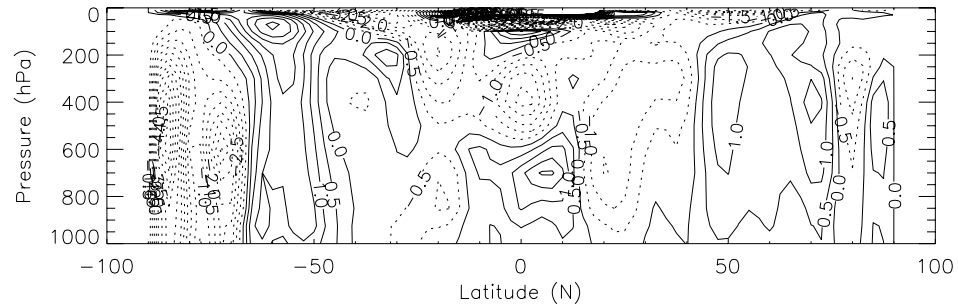


(Haigh, *Phil.Trans.Roy.Soc.* 2003)

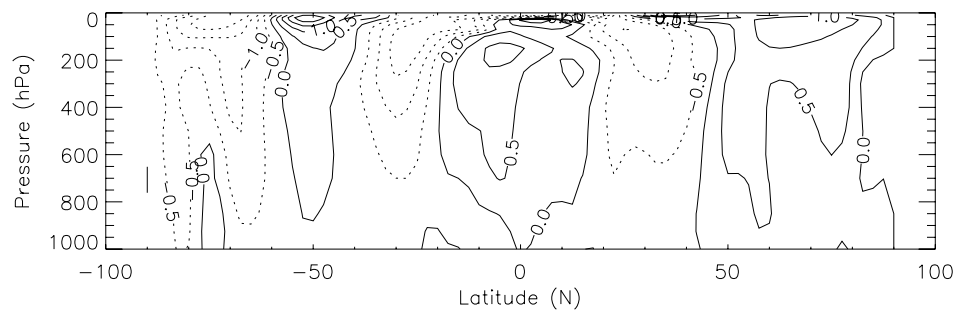


mean

# Influences on zonal wind



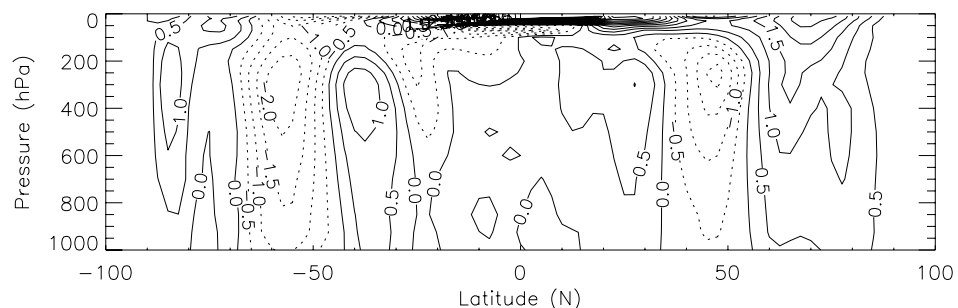
trend



solar

***Jets weaken & move polewards***

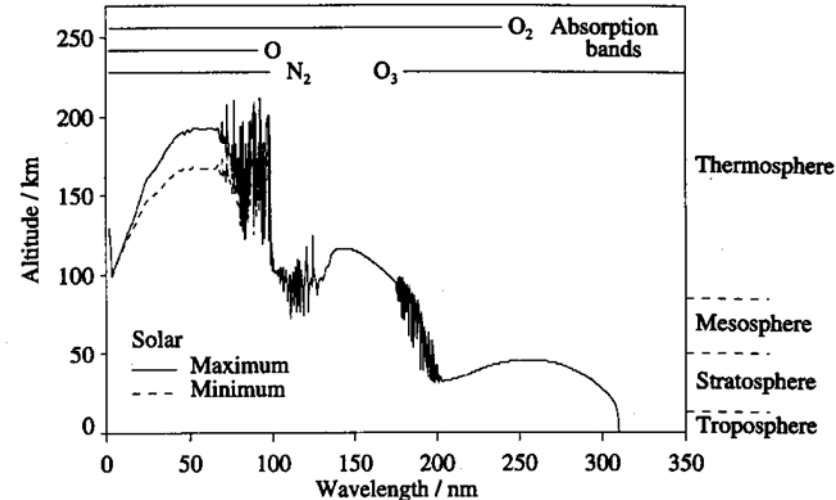
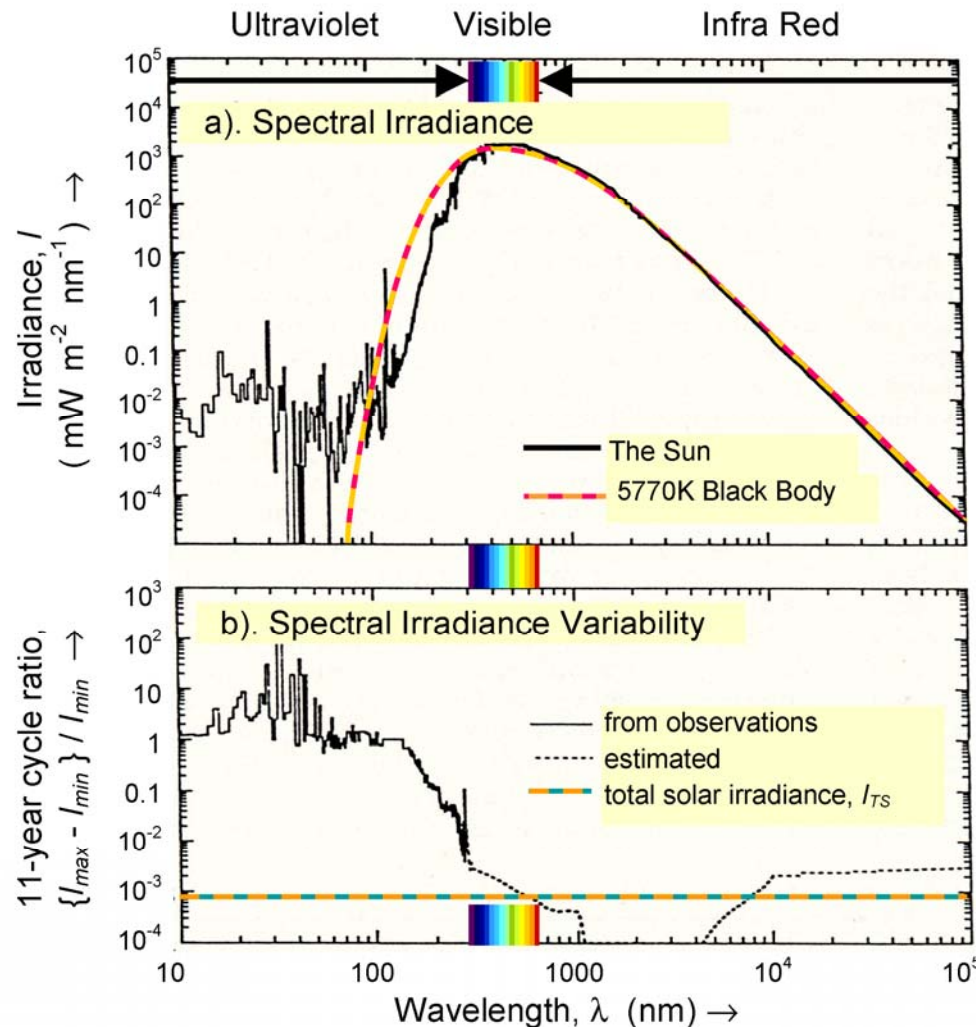
***Can we explain this?***



volcanic

(Haigh et al, *J.Clim.*, 2005)

# Solar spectrum 11-year cycle variability



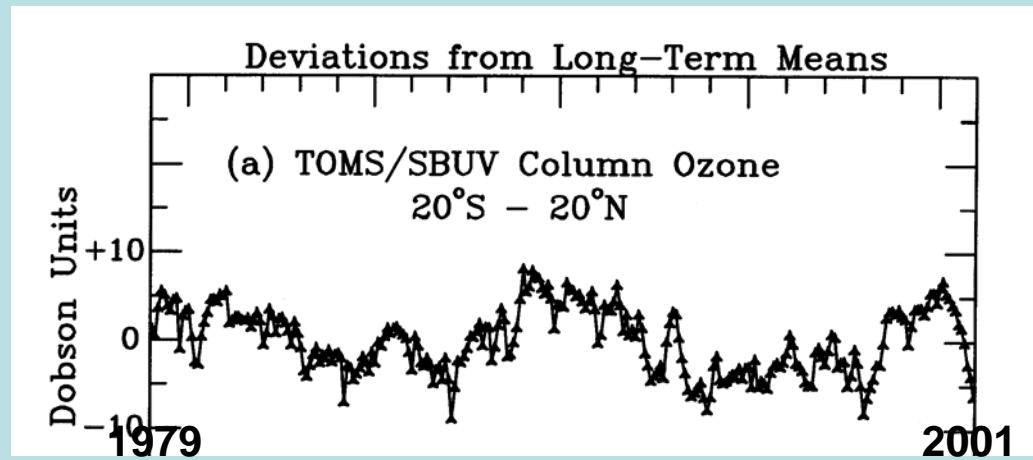
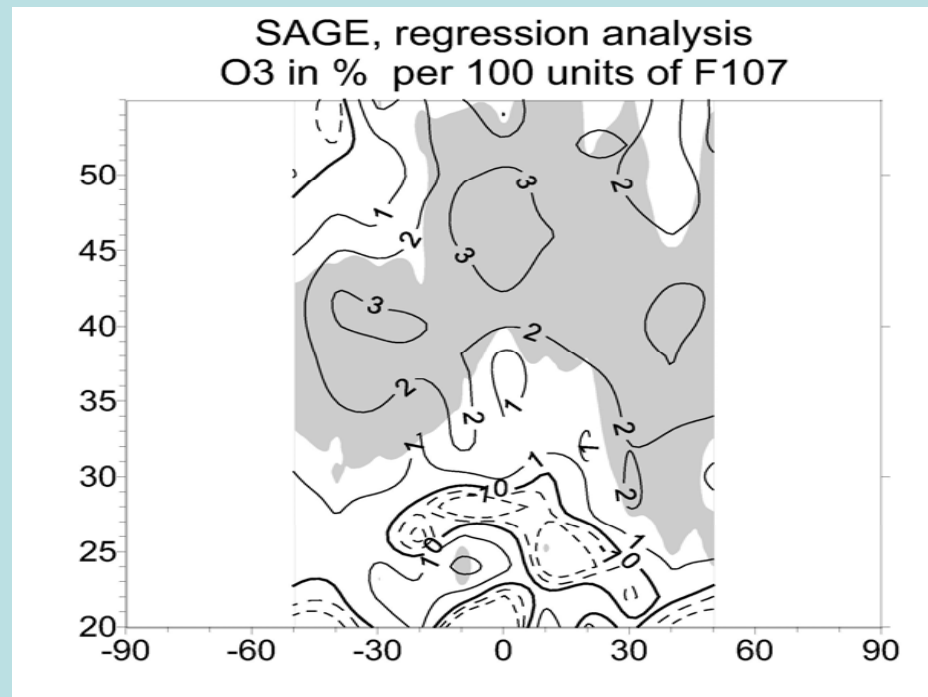
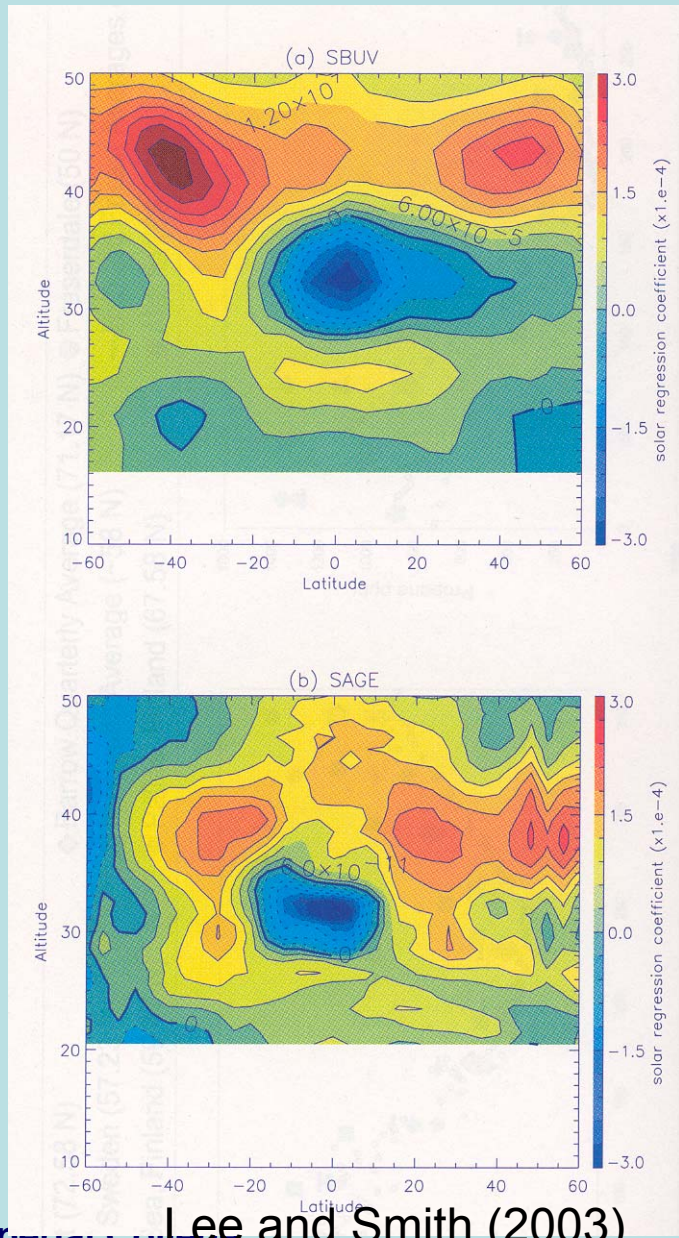
Altitude of unit optical depth

Andrews(2000)

***Larger changes in UV  
tend produce ozone and  
heat the stratosphere***

Lean (1991) ed. Lockwood.

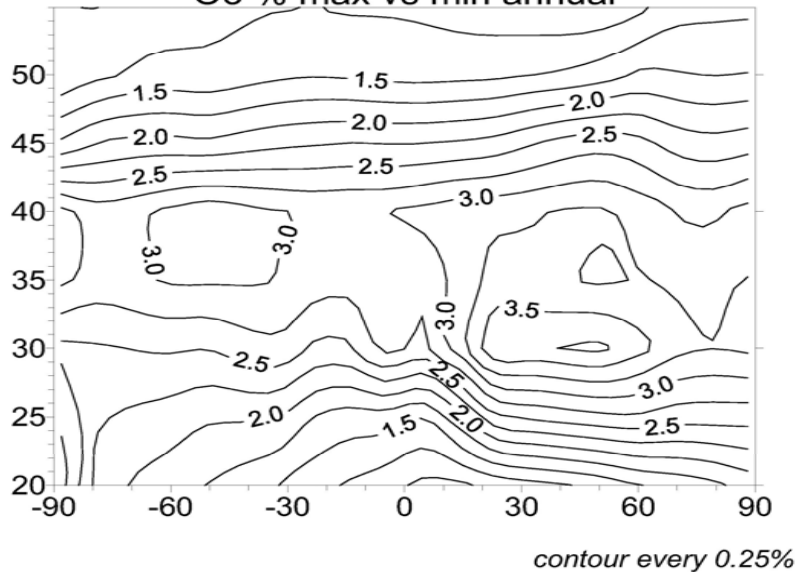
# Ozone (observations)



Hood and Soukharev (2003)

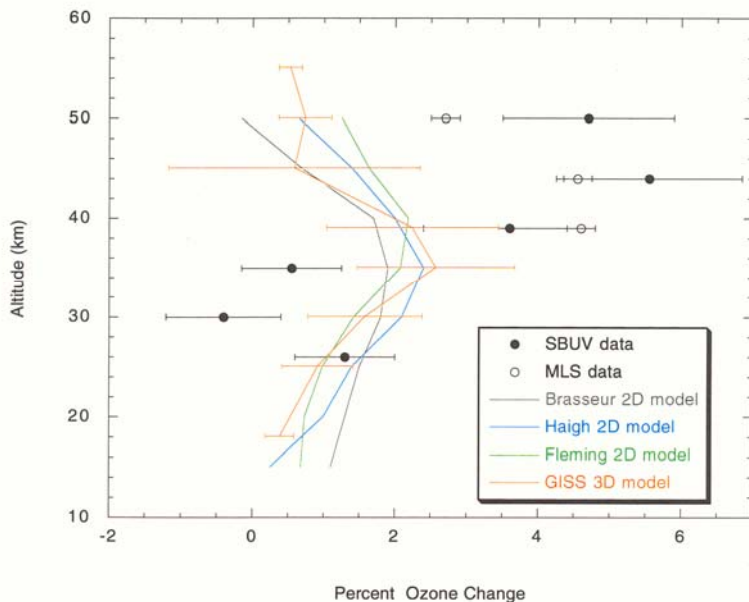
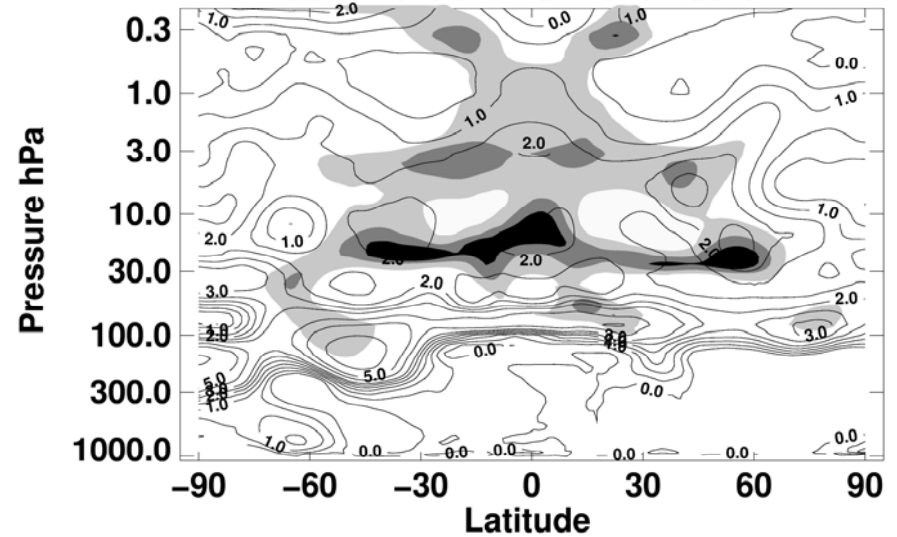
# Ozone (models)

OSLO MODEL  
O3 % max vs min annual

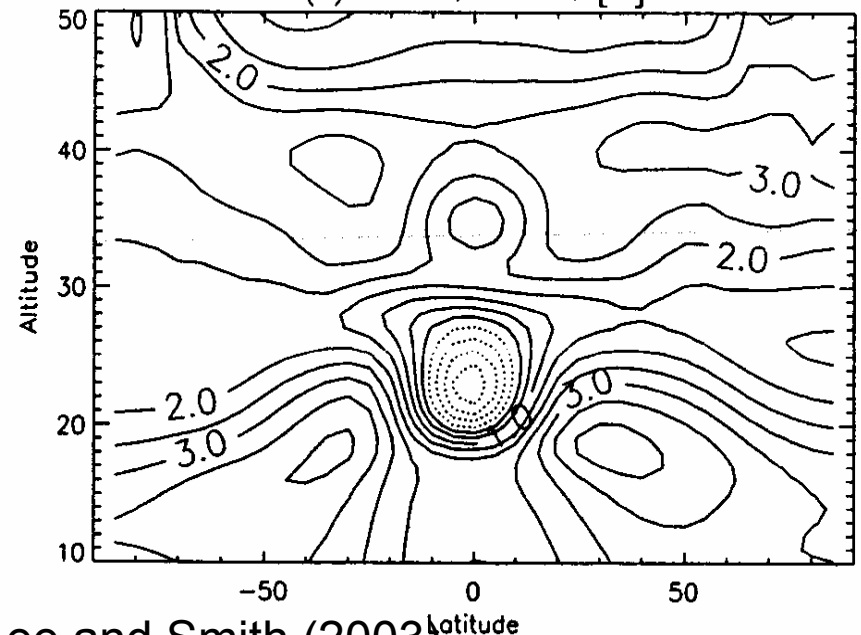


UKMO

Ozone solar difference %



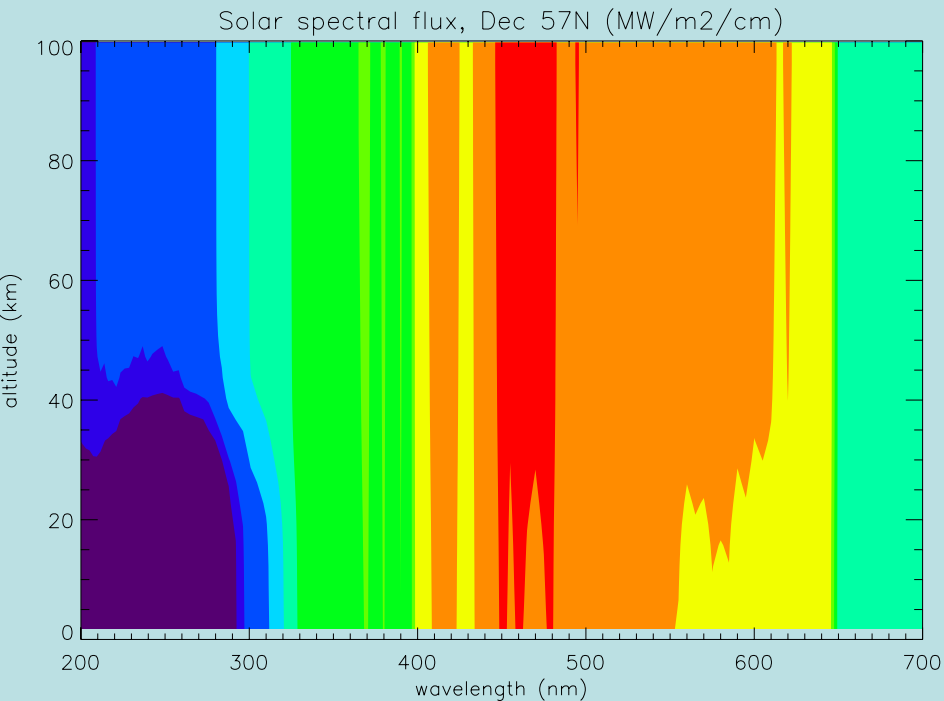
(c) SOL+QBO+VOL [%]



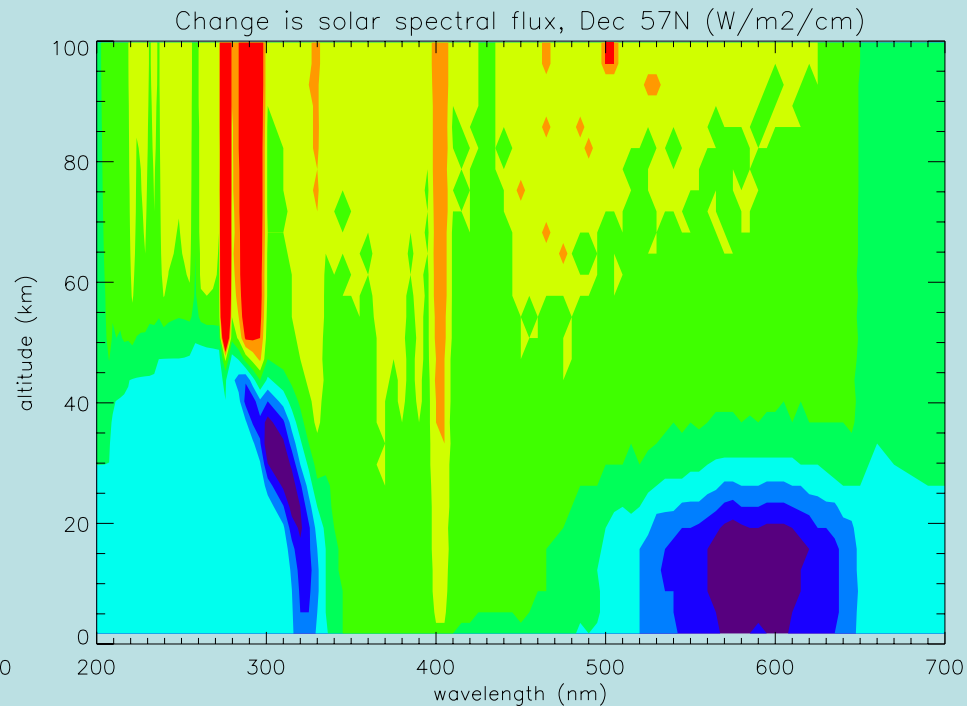
Lee and Smith (2003)

# Spectral irradiance: how ozone impacts radiative forcing

solar min



max-min



(Haigh, *Nature*, 1994)

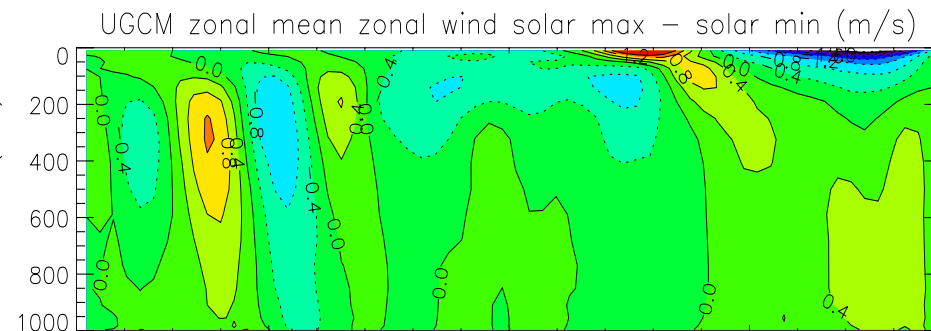
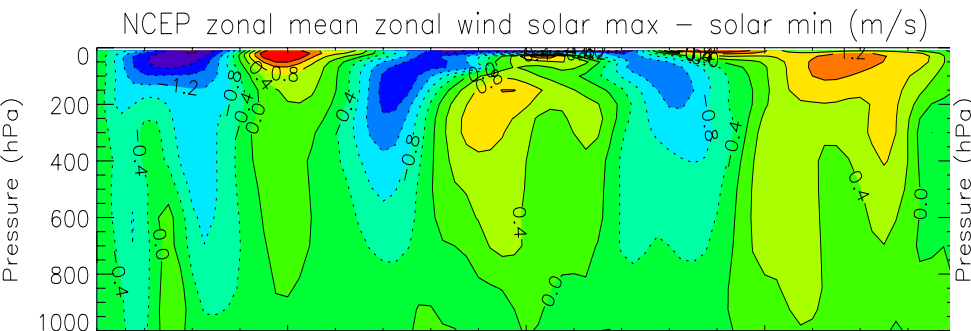
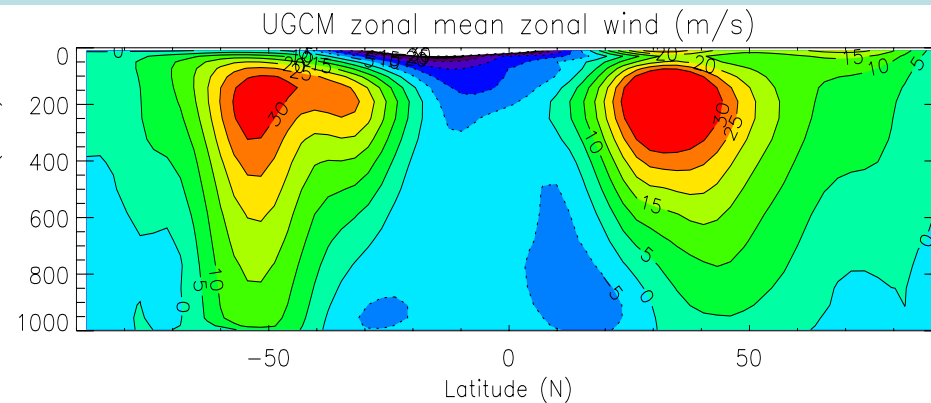
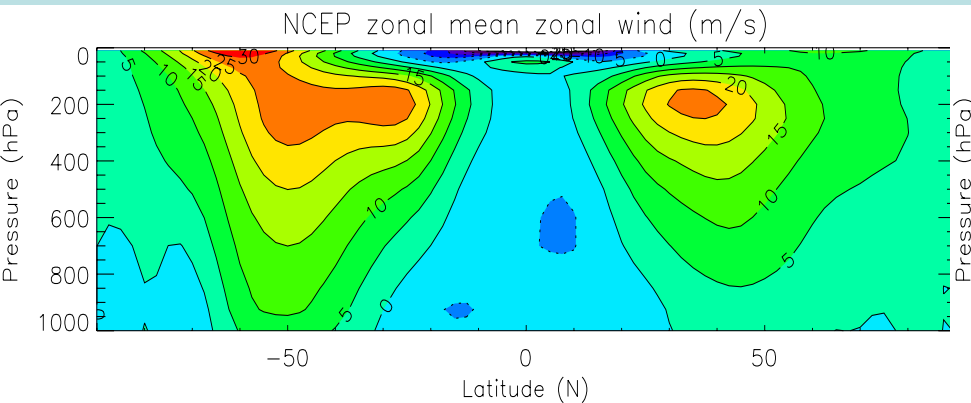
# Solar influence on zonal mean zonal wind

**Observations**

**Model with enhanced UV**

- *not bad!*

- *BUT can we explain this pattern?*



Haigh et al, *J.Clim.*, 2005

Haigh, *Science*, 1996

# Potential mechanisms

- Increased static stability weakens tropical upwelling and Hadley cells
- Changes in  $u$  affect growth/propagation of planetary waves
- Baroclinic lifecycles (mid-latitude weather patterns) modified

# Simplified GCM - “dynamical core” model

Based on University of Reading spectral model:

Full dynamics T42 L20.

No orography.

Newtonian cooling - equinoctal radiative equilibrium temperatures  $T_e(\text{lat.}, \text{ht.})$ .

Rayleigh friction.

Experiments:

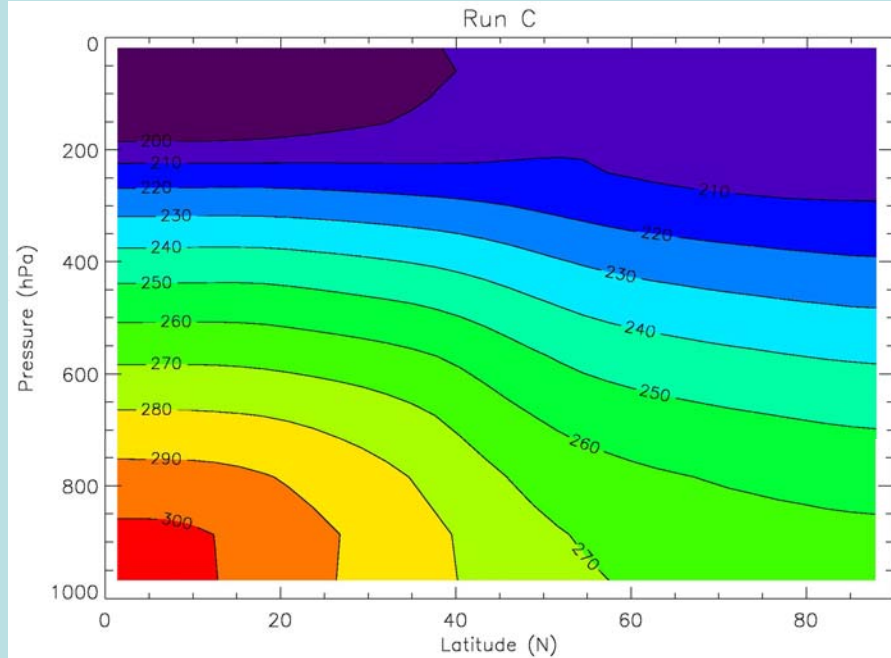
1. Equilibrium response to perturbations to stratospheric  $T_e$ .  
(Haigh, Blackburn & Day, *J.Clim.*, 2005)
2. Spin-up ensemble.

# Simplified GCM perturbation experiments

- Run C**      Control:  $T_e$  distribution of Held and Suarez (1994).
- Run U5**      Stratospheric **only**  $T_e$  increased uniformly by 5K.
- Run E5**      Stratospheric **only**  $T_e$  increased by 5K at the equator, decreasing with  $\cos^2(\text{latitude})$  to 0K at the poles.

Simplified GCM  
perturbation runs:

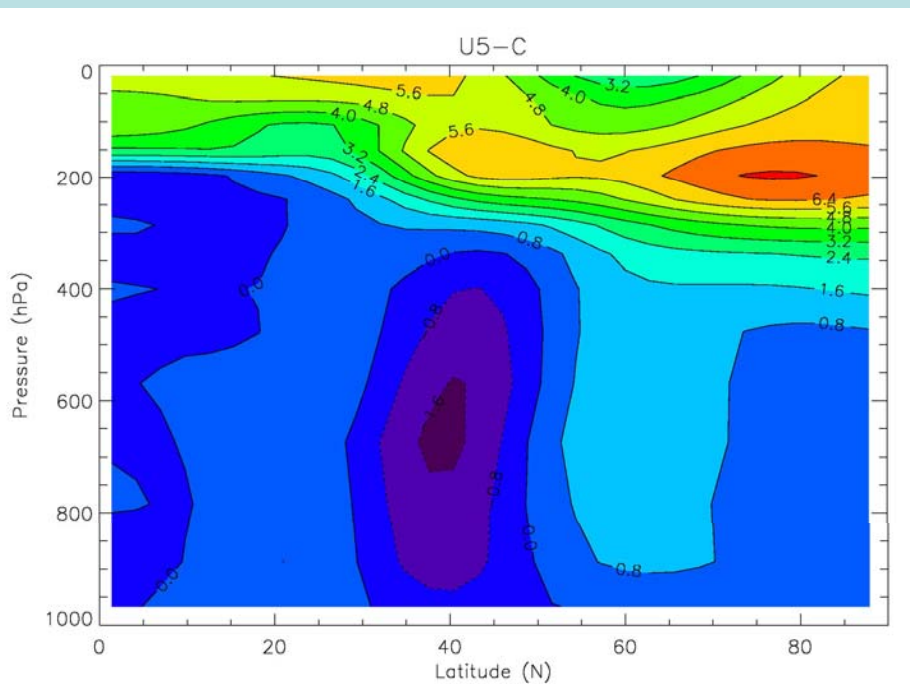
temperature (K)



Run C

U5 - C

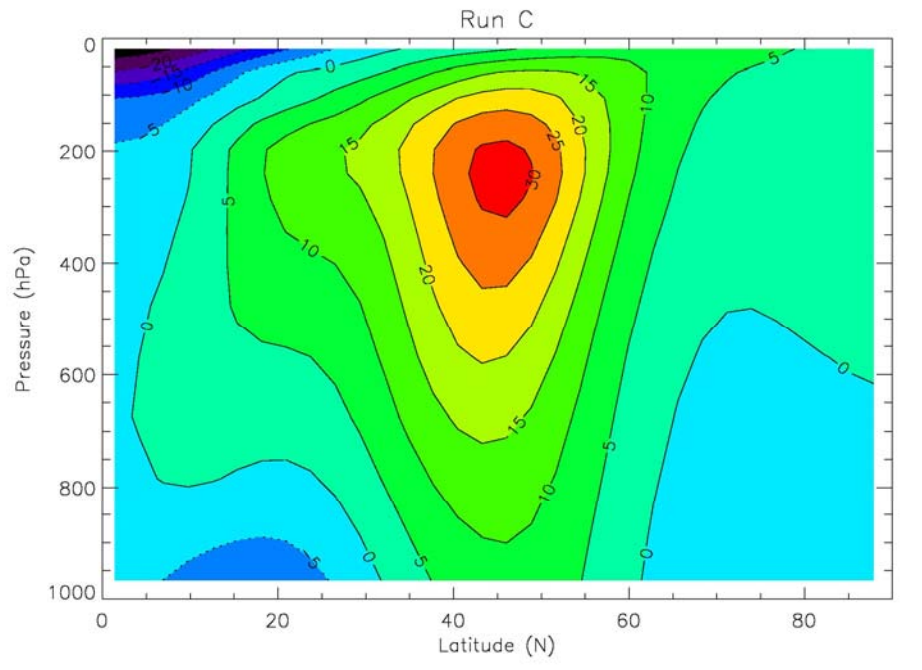
E5 - C



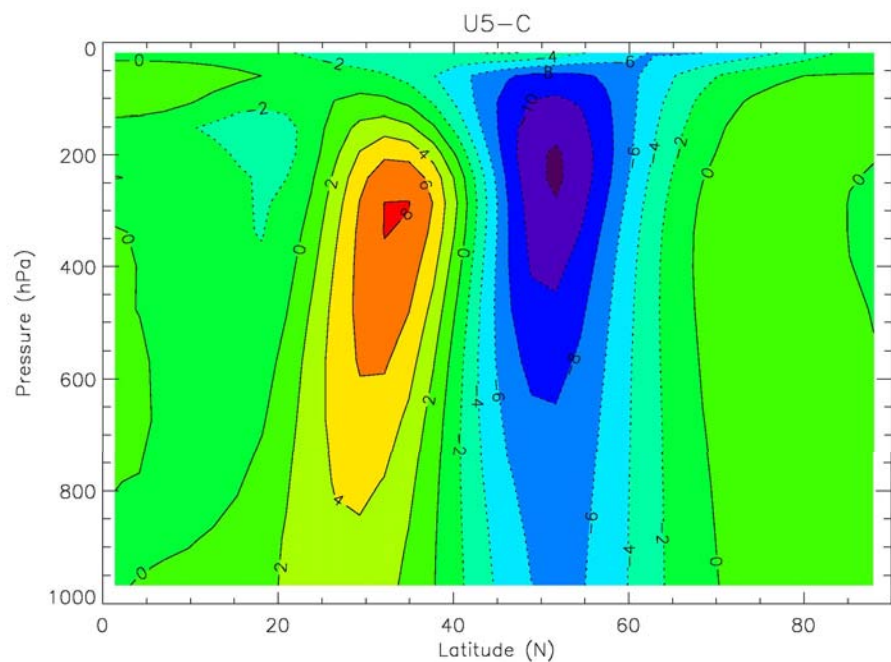
# Simplified GCM perturbation runs:

zonal wind (m/s)

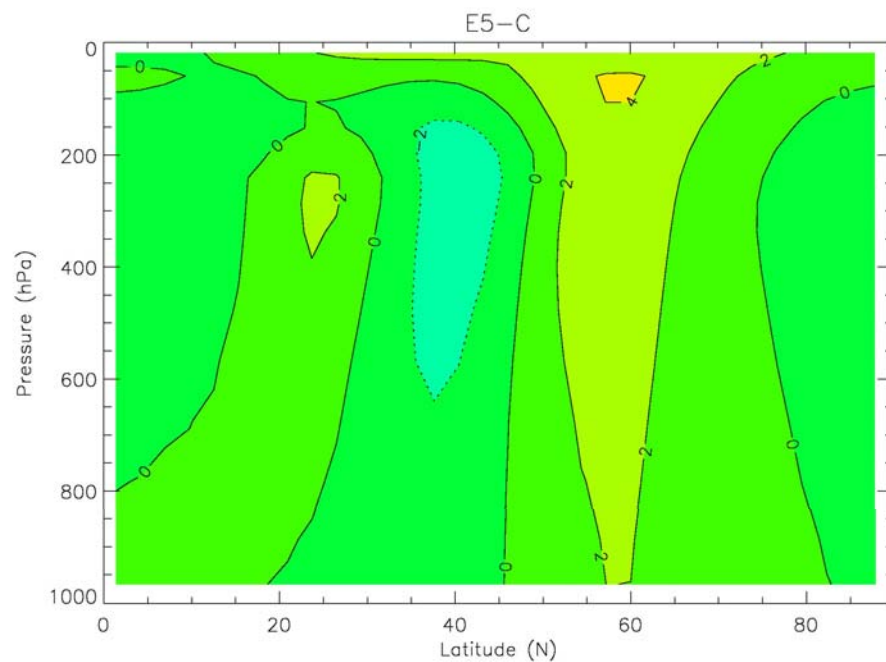
Run C



U5 - C



E5 - C



# Summary

- Detectable signals of solar influence on climate (on many timescales)
- Non-uniform patterns of warming
- Standard climate models cannot explain this
- Enhanced solar UV heating in the stratosphere may provide a clue
- – but mechanisms are not yet fully understood

